Agronomy 406 World Climates

February 1, 2018

Ocean properties and circulation.

Measuring the oceans.

Team 2 Climate News on Tuesday.

Review for today:

Rahmstorf, S.: The Ocean Thermohaline Circulation: A Brief Fact Sheet

For Tuesday:

Online textbook: 5.2.1 El Niño-Southern Oscillation El Niño, La Niña, and ENSO FAQ

Climate News

See notes for January 18 regarding things to focus on when preparing your Climate News story.

Do background work to make sure you understand the original source and the basic facts relevant your story (terms, locations, etc).

Your team's presentation should tell a story with a beginning, middle, and end rather than separate pieces stuck together.

Also read "Recommendations for presentations" in the Pages section on the AGRON 406 Canvas page.

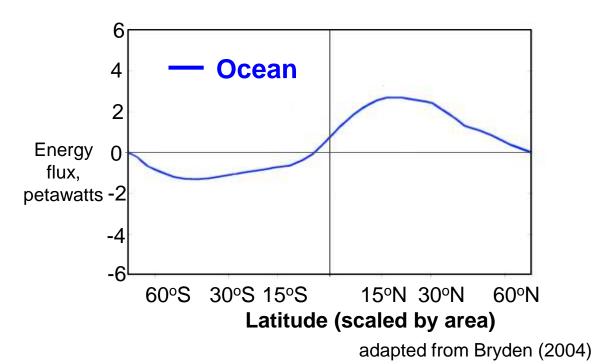
Bring your presentation on a USB stick and load it before the start of class.

Ocean circulations

Ocean circulations transport energy comparable to the atmospheric sensible or latent heat fluxes.

This has big effects on climate, such as warming by the Gulf Stream and North Atlantic Drift.

Look at processes that cause ocean circulation.

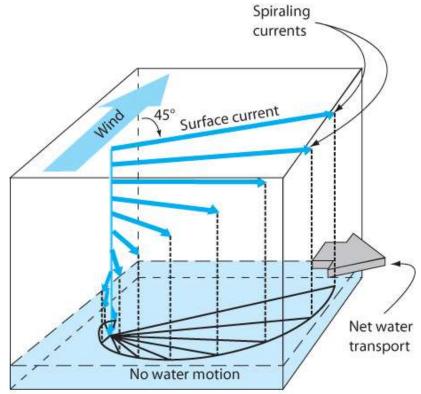


Water moves to the right of the wind

Each deeper layer moves a little more to the right than the layer above.

Looking from the top, the motion appears to "spiral" rightward with depth.

- This is the Ekman spiral (after V. Ekman, who first solved the equations for it).
- Averaged over the whole depth, the water moves 90° to the right of the wind.



(a) Ekman spiral in the northern hemisphere

- Theoretical result - real world differs somewhat.

Wind forcing causes convergence of ocean water in subtropics

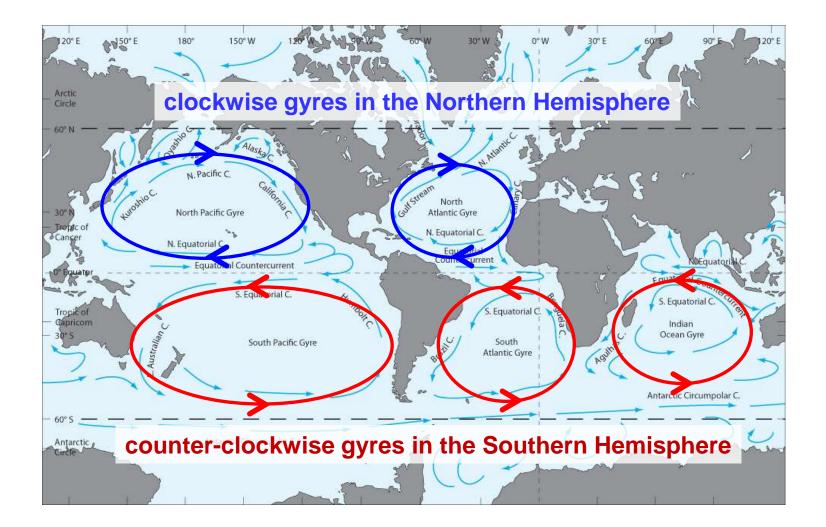
Given:

- Winds in the mid-latitudes are westerly.
- Winds in the tropics are easterly.
- Coriolis force causes water to move to the right of the wind in the Northern Hemisphere.



This causes water to converge in the region between the easterlies and westerlies.

Ocean gyres



Temperature and salt

Ocean water is affected by **salt content:**

– Average salinity is about 35 parts / thousand ($^{\circ}/_{\circ\circ}$).

Salty water freezes at lower temperature than fresh water.

Density depends on both temperature and salinity:

- Density increases with increasing salt content.
- Density increases with decreasing temperature.
- Unlike fresh water, sea water becomes denser with colder temperature until it freezes.

What processes in the climate system can change the **temperature** of ocean water?

What processes can change its **salinity**?

Discuss in your teams:

Think of two processes that can change the salinity of the ocean.

Think of two processes that can change the **temperature** of the ocean.

Processes that can change density

Change temperature

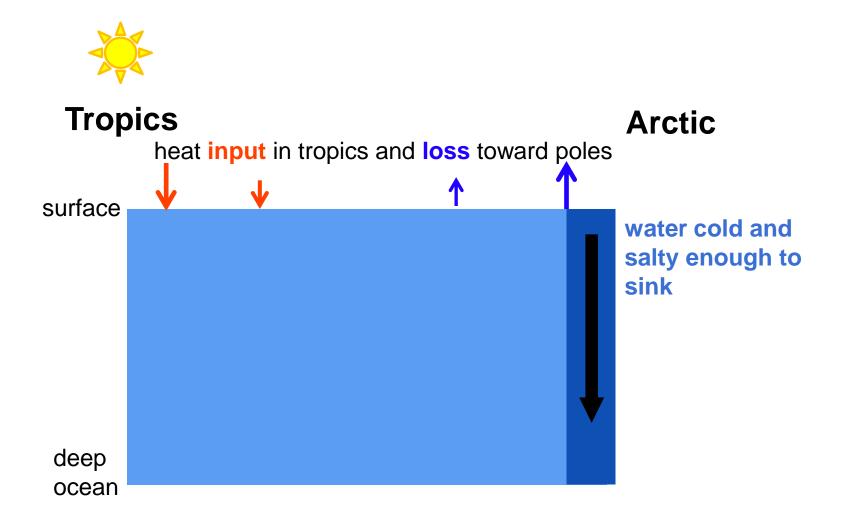
- ice melting
- cloud cover
- cold front
- daytime heating
- evaporation (latent heat)
- greenhouse effect (global warming)
- sea vents
- upwelling and downwelling

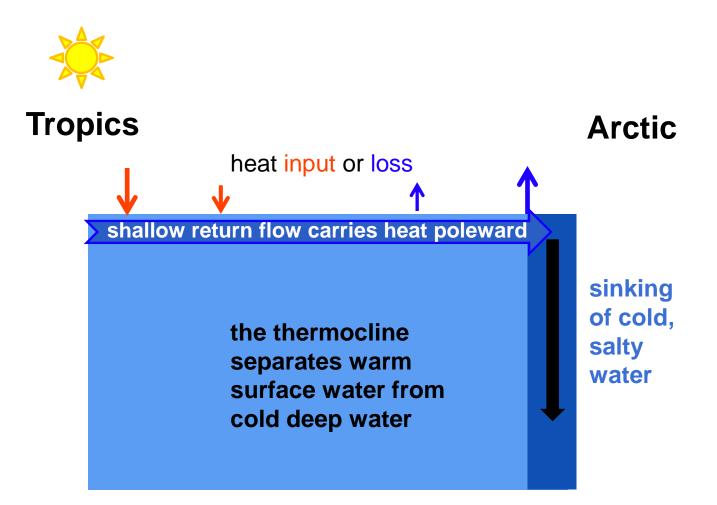
Change salinity

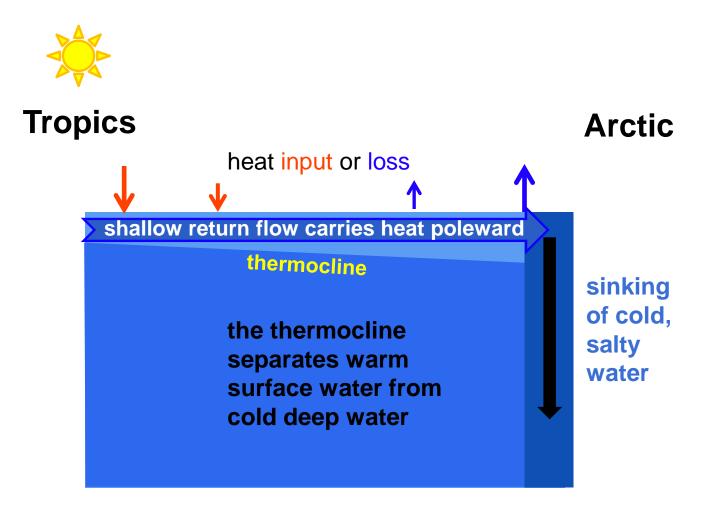
- ice melting
- precipitation
- fresh water from rivers
- evaporation
- circulation in the ocean
- upwelling and downwelling

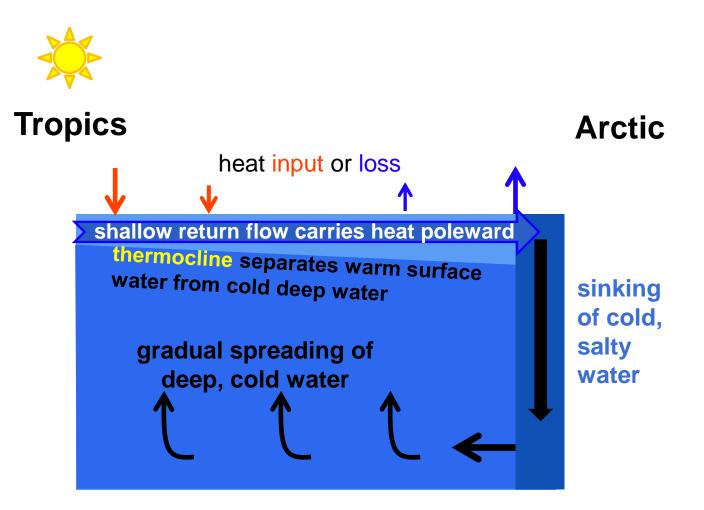
When you open the drain in a bathtub, what happens to water at the other end of the tub?



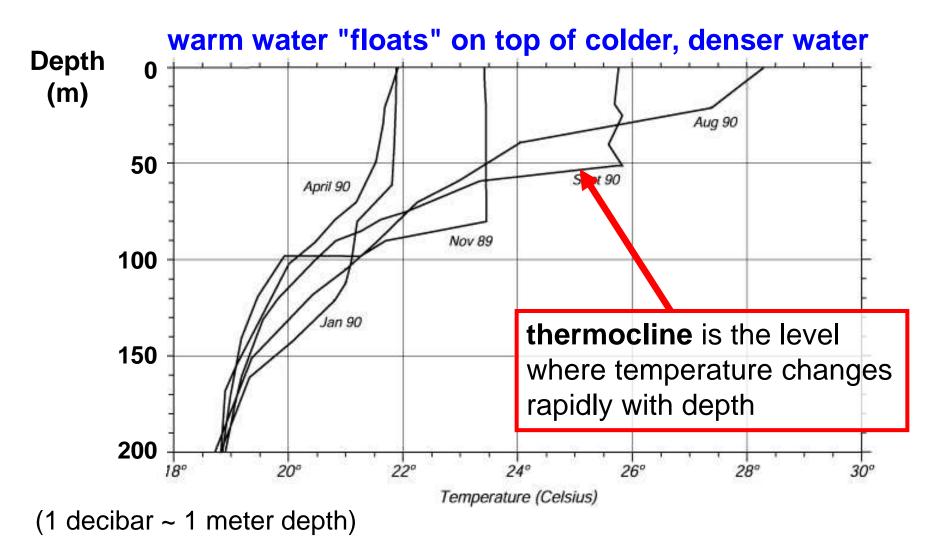




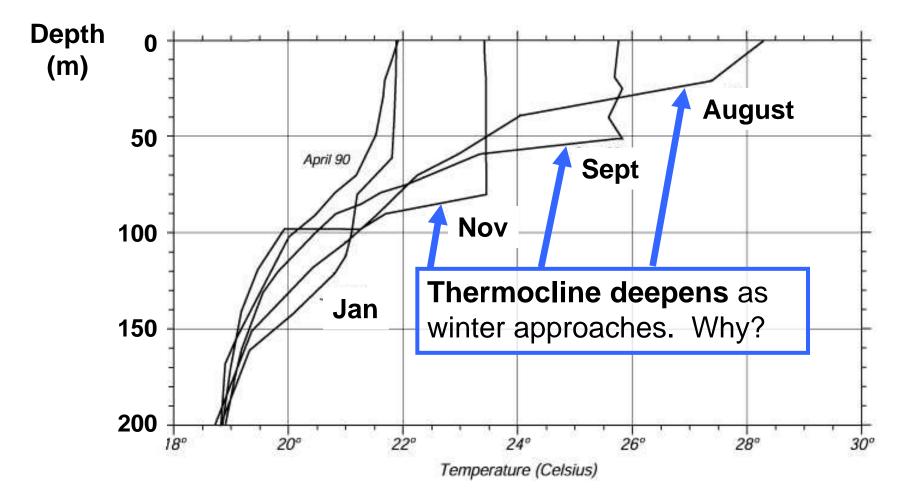




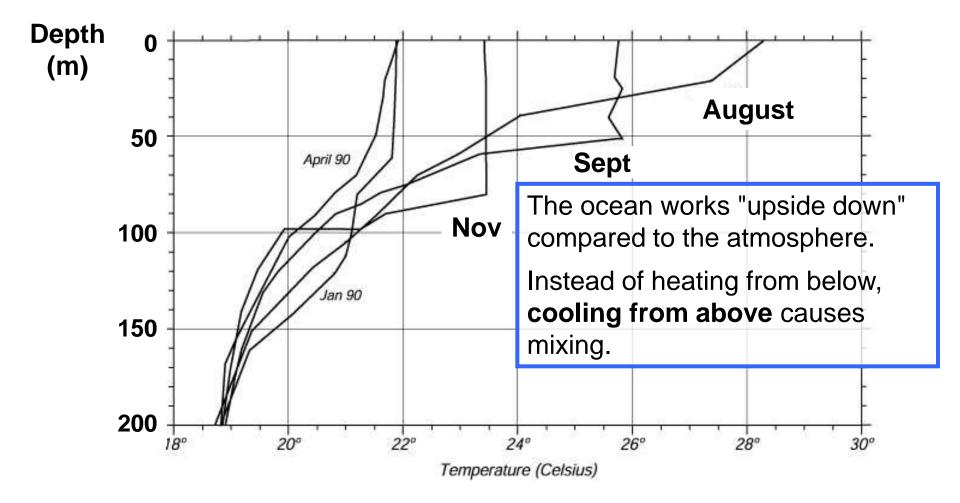
Vertical temperature profile of the ocean



Vertical temperature profile of the ocean



Vertical temperature profile of the ocean



Thermohaline circulations: Deep water formation

Sinking of cold, salty water in **deep water formation** regions is the main control on the thermohaline circulation.

There are only a few deep water formation regions: Greenland-Norwegian Seas (north of Iceland) Labrador Sea (between Canada and Greenland) Weddell Sea (off Antarctica, on the Atlantic side) Ross Sea (off Antarctica, on the Pacific side) Mediterranean Sea (off the south coast of France)

Thermohaline circulation: The ocean conveyor belt

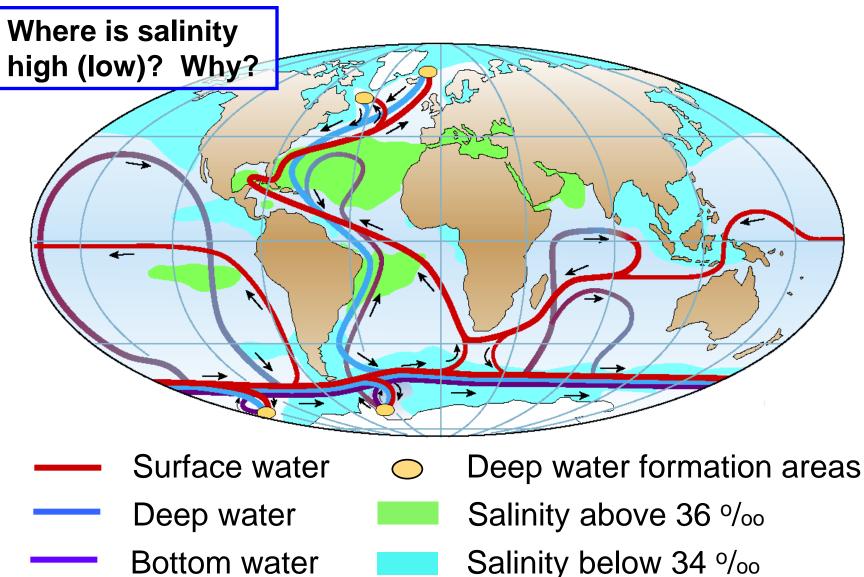
Sinking of cold, salty water in deep water formation regions is the main control on the **thermohaline circulation**.

Combined with wind-driven circulation, produces the **ocean conveyor belt** that connects surface and deep waters.

Thermohaline circulations move slowly: Time to make a complete transit around the ocean conveyor is of order **1000 years.**

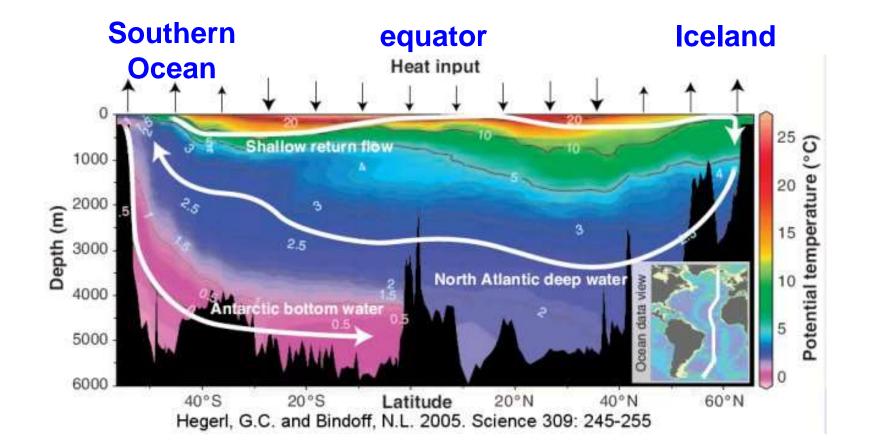
The conveyor transports energy (heat) around the world. Also transports whatever is dissolved in the ocean – salt, oxygen, nutrients.

The ocean conveyor belt



from Rahmstorf (2002), Nature

Actual thermohaline circulation is a little more complicated



more ocean cross sections available at http://www.ewoce.org/gallery/index.html

Why do we care?

The ocean conveyor belt transports heat poleward – important to the global energy balance.

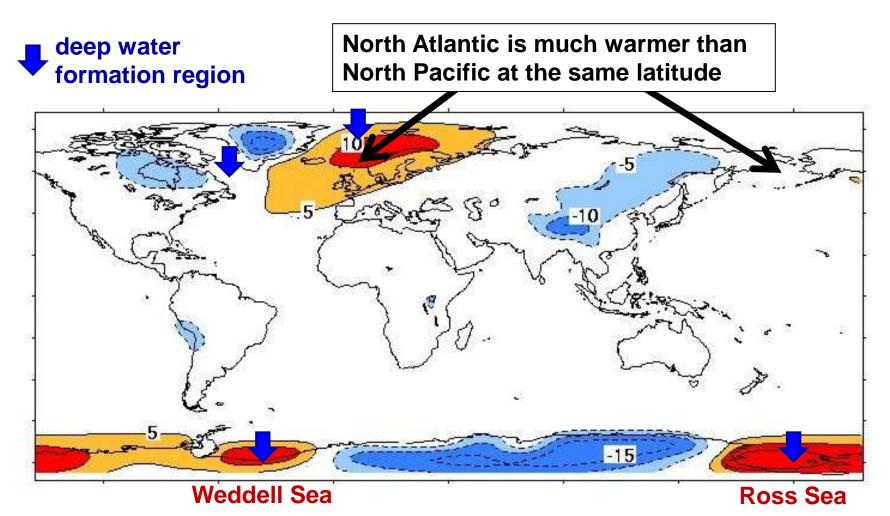
Sinking carries heat into the ocean:

- Most of the excess heat due to increasing greenhouse gases has gone into ocean, down to about 700 meters.
- But probably not the "very deep" or abyssal part of the ocean (deeper than ~2000 meters).

Poleward heat transport in the ocean conveyor belt keeps some regions much warmer than they would be otherwise.

Thermohaline circulation has important effects on temperature, especially in the North Atlantic

Temperatures shown as deviations from the average at each latitude



Ocean measurements

- Satellites
- Buoys
- Ships

Each has advantages and disadvantages.

Satellite observations

Sea surface temperature:

Measured from infrared and microwave radiometers.

Corrections have to be applied for effects of clouds, atmosphere and solar radiation.

Only "sees" the ocean surface.

Winds and surface stress:

A **scatterometer** works sort of like a radar. Waves on the surface affect the amount of energy that is scattered back to the satellite.

Measurements of the same point from different angles allow estimates of wind speed and direction.

Salinity

SMOS (Soil Moisture and Ocean Salinity) SMAP (Soil Moisture Active and Passive) Aquarius

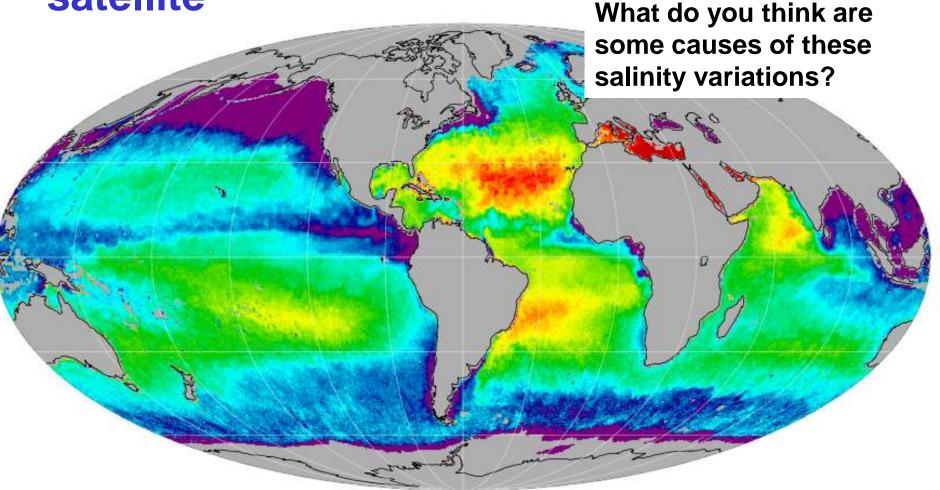
Salinity affects electromagnetic properties of water.

The satellite measures electromagnetic radiation from the sea surface, which is then converted to sea surface salinity.

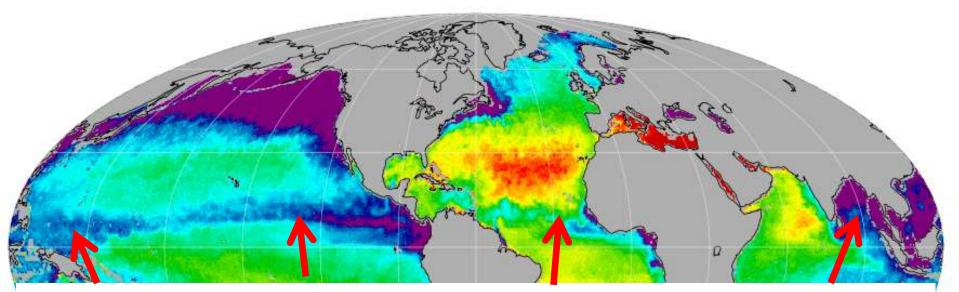
Only "sees" the ocean surface.



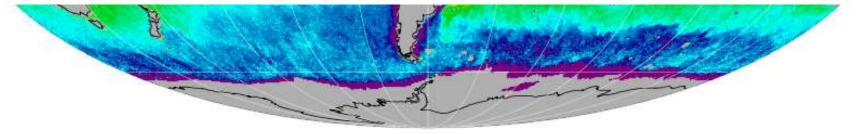
Most recent sea surface salinity from the Soil Moisture Active Passive satellite What do you thin



Some reasons for salinity variations



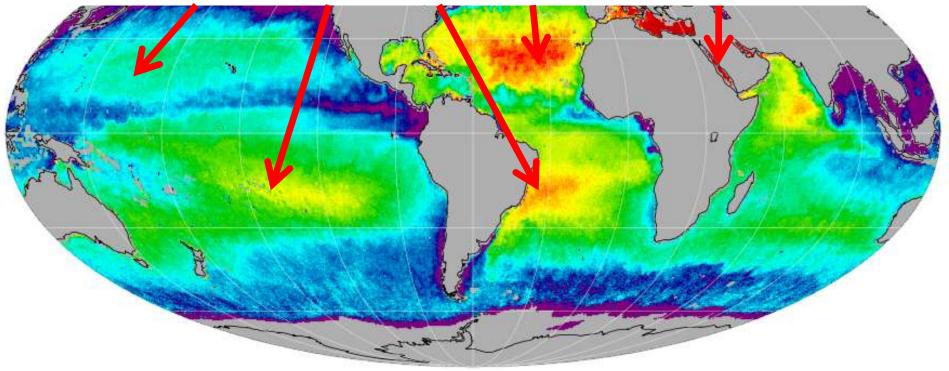
Average position of the rising branch of Hadley cell (ITCZ): cloudiness reduces evaporation, and precipitation adds fresh water



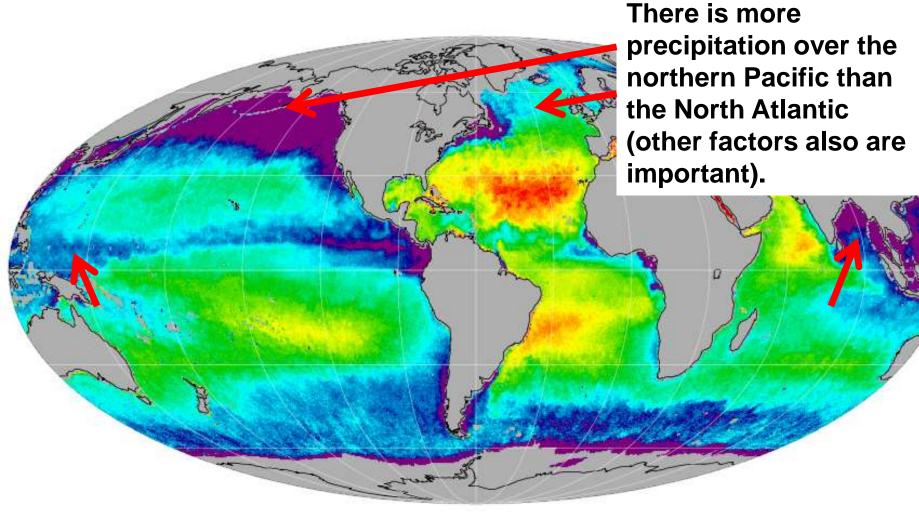
Some reasons for salinity variations



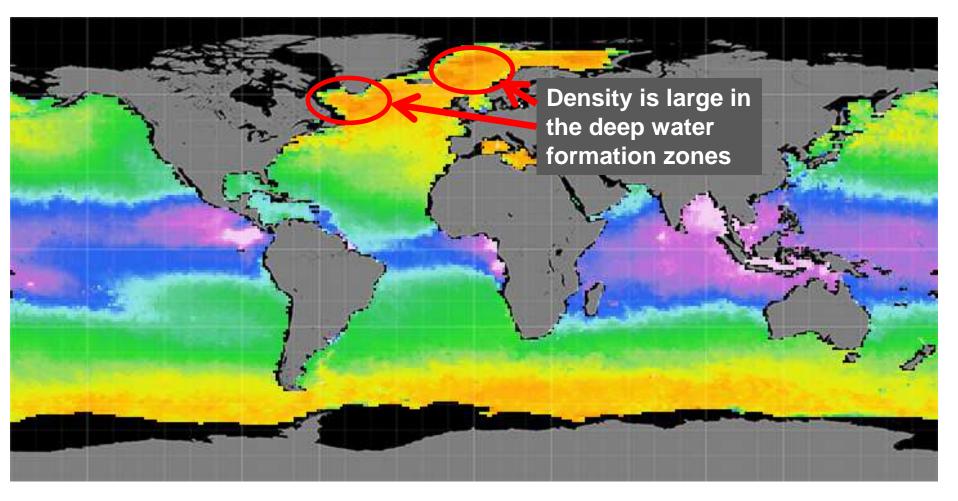
Sinking branches of Hadley cells: clear skies increase evaporation, leaving salt behind



Some reasons for salinity variations



Sea surface density from the Aquarius satellite



April 2015

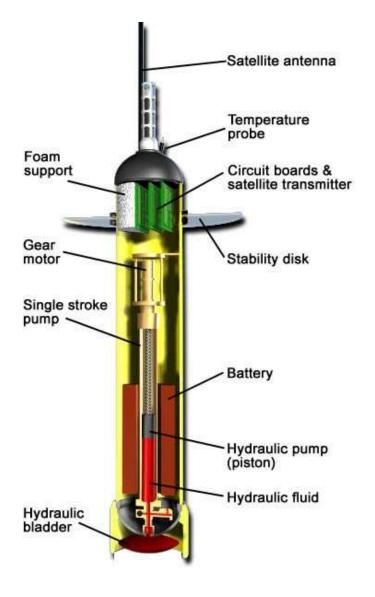
Always remember...

- Satellites **do not** measure temperature.
- Satellites **do not** measure wind speed.
- Satellites do not measure salinity.

Satellites measure **amount of radiative energy flux** (radiance) at certain wavelengths.

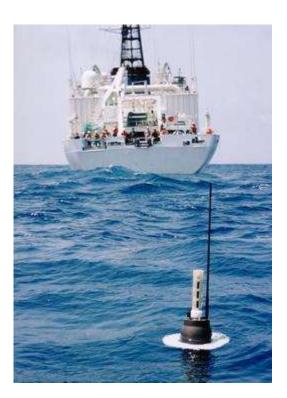
Radiances have to be converted to the variables we want using mathematical algorithms and corrections. These conversions can be uncertain and prone to error.

Argo floats



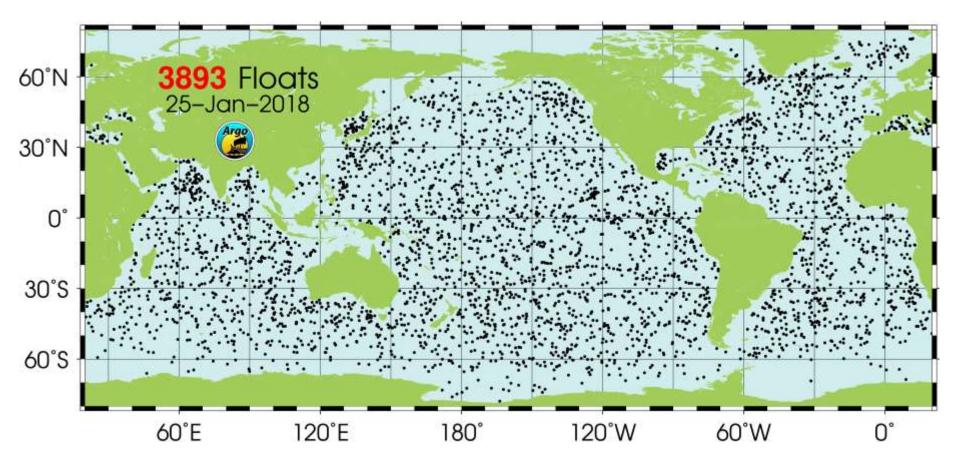
Spends most of its time drifting at a depth of 1000 m.

About every 10 days, the float descends to 2000 m and takes measurements as it rises to the surface.



Video of float operation

Positions of Argo floats that reported data in the 30 days before January 25, 2018



from http://www.argo.ucsd.edu/

Argo floats

Floats don't last forever:

A float can go through about 150-200 cycles of ascent and descent before its battery gets too weak.





About 800 new floats must be deployed per year to maintain the current network.

Operating one float costs about \$30,000 (about half for the float itself, and half for data processing and infrastructure).